Numerical study of droplet growth by condensation by means of massive DNS simulations



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Rain formation



- Activation of Cloud Condensation Nuclei (CCN)
- Growth by condensation
- Growth by mixing/collision due to turbulence
- Growth by gravitational collision



Turbulence and droplet Condensation DNS: a Brief Excursus Sto



- turbulence has been indicated as the key missing link to solve condensation/ collision coalescence problem
- first DNS of turbulence/cloud interactions done by Vaillancourt et al. 2002.
 Domain 10cm, resolution 80³ grid points, droplets 50000→ Conclusion:
 negligible effect of the small-scale turbulence on droplet spectra broadening
- Celani et al. 2007, resolving large-scale fluctuations 2D cloud → Conclusion: dramatic increase in the width of the droplet spectrum is qualitatively found although the dynamics of the small scales is not resolved
- Paolo & Sharif, 2009, same conclusions but 3D simulation obtained adding an arbitrary large-scale forcing on the supersaturation equation field.
- Current state of the art: Lanotte et al 2009, 3D DNS simulations increasing size of the cloud up to 70 cm→ turbulence affects droplet spectra broadening mechanism by increasing the cloud size.



Lanotte, Seminara, Toschi, JAS 2009





- Reynolds number : Importance of large scales
- Upper limit $T_L < au_s$ $\sigma_{R^2} \propto A_3 A_1 v_\eta au_\eta^2 R e_\lambda^{5/2}$
- Lower limit $T_L > \tau_s$ $\sigma_{R^2} \propto A_3 A_1 v_\eta \tau_\eta \tau_s R e_\lambda^{3/2}$



Our objectives



- 1) Has Droplet spectra variance in warm cloud been well approximated so far?
- 2) Metodologies: -Direct Numerical Simulation DNS
- Current simulation time seconds/2 minutes→ up to 20 minutes

Turbulence and condensation: mathematical model



Eulerian framework: Navier-Stokes + supersaturation field s \rightarrow s>0 condensation s<0 evaporation

→ s>0 condensation s<0 evaporation $\partial_t \mathbf{u} + \mathbf{u} \cdot \nabla \mathbf{u} = -\frac{1}{\rho} \nabla p + \nu \nabla^2 \mathbf{u} + \mathbf{f}, \quad \nabla \cdot \mathbf{u} = 0$

$$\partial_t s + \mathbf{u} \cdot \nabla s = \kappa \nabla^2 s + A_1 w - \frac{s}{\tau_s}$$

Possible large scale forcing of supersaturation

$$\tau_s^{-1} = \frac{4\pi\rho_w A_2 A_3}{V} \sum_{i=1}^N R_i$$

Phase relaxation time scale

 $\begin{aligned} & \frac{d\mathbf{V}_{i}(t)}{dt} = -\frac{\mathbf{V}_{i} - \mathbf{v}_{i}[\mathbf{X}_{i}(t), t]}{\tau_{d}} + g\mathbf{z} & \text{pressure} \\ & \frac{d\mathbf{X}_{i}(t)}{dt} = \mathbf{V}_{i}(t) & \text{Formula} \\ & \frac{dR_{i}(t)}{dt} = A_{3}\frac{s[\mathbf{X}_{i}(t), t]}{R_{i}(t)} & \text{Same formula} \end{aligned}$

Droplet modeled as point particles

Force acting on droplets: Stokes drag and gravity

Same formulation of Lanotte et al., JAS 2009

Numerical Methodology



- Combined Eulerian/Lagrangian Solver
- Pseudo-spectral code
- 2/3 rule for dealiasing
- Tri-linear interpolation to evaluate fluid velocity and saturation field at the droplet position
- Tri-linear extrapolation to calculate droplet feedback on the saturation field
- Full MPI parallelization for both carrier and dispersed phase
- Computational time step linearly scales up to 10000 cores → huge simulations

Simulations parameters

									North
Label	N^3	L_{box}	v_{rms}	T_L	T_0	Re_{λ}	N_{c}	ł	Stockholme
		[m]	[m/s]	[s]	[s]				JUDICKHOITHS
DNS $A1/2$	64^{3}	0.08	0.035	2.3	0.64	45	6×1	10^4	universitet
DNS $B1/2$	128^{3}	0.2	0.05	4	0.95	95	$9.8 \times$	10^{5}	
DNS $C1/2$	256^{3}	0.4	0.066	6	1.5	150	9×1	10^6	
DNS D1	1024^{3}	1.5	0.11	14	3	390	$4.4 \times$	10^{8}	
DNS E1	2048^{3}	3	0.12	30	4	600	3. imes	10^{9}	
LES E1	512^{3}	100	0.7	142	33	5000	$1.3 \times$	10^{14}	
Dissipation r Kolmogorov	ate: ٤ scale:	$arepsilon = \eta$	= 10	-3 [m	$\frac{n^2s}{m}$	3^{-3}	Typi s	cal value found in tratocumuli
	11	—		1		2			

- Kolmogorov time: $au_\eta = 0.1$ s
- Initial Radius: 13 µm (1) 5 µm (2) $St_{\eta} = 3.5E 2 \div 5E 3$ Phase relaxation time: 2.5 s (1) 7 s (2) $C = 130/cm^3$

DNS E1: state of the art with 1024³ grid point resolution corresponding to a domain length of 1.5 meters with 10⁹ droplets evolved. First DNS with cloud size order meter.

E1, C2 and D1 do not reach 20 minutes of simulation

Conservative hypothesis



- We assume $<s>=0 \rightarrow$ all is due to s fluctuations
- No mean updraft
- Consequently $\langle R^2 \rangle = R_0^2$
- Entrainment effects are not considered (Kumar, Schumacher and Shaw, JAS 2014)
- Adiabatic approximation → small temperature fluctuations (DNS D1 T_rms=0.005 K) → A₁, A₂, A₃ costants
- Effects due to inhomogeneity are not captured
- Collisions not included
- Our results represent a lower limit on droplet growth



 $\frac{d\langle (R_i^{2'})$

dt

DNIC maguilta



- σ_{R^2} standard deviation of square radius fluctuations
- Standard deviation increases continuosly even if s has reached the quasi steady state
- Power law t^{1/2}
- Proportional to $\operatorname{Re}_{\lambda}$ and \mathcal{T}_{S}
- Larger scales are responsible for variance growth
- Correlation $\langle s' R^{2'} \rangle$ reaches a quasi-steady state

$$\frac{\left| {}^{2} \right\rangle}{dt} = \frac{d\sigma_{R^{2}}^{2}}{dt} = 4A_{3} \langle s' R^{2'} \rangle$$

$$\begin{aligned} \mathbf{1-D \ stochastic \ model/1} \\ w_i'(t+dt) &= w_i'(t) - \frac{w_i'(t)}{T_0} dt + v_{rms} \sqrt{2 \frac{dt}{T_0}} \xi_i(t)^{\text{Stockholms}} \\ s_i'(t+dt) &= s_i'(t) - \frac{s_i'}{T_0} dt + A_1 w_i' dt - \frac{s_i'}{\langle \tau_s \rangle} dt + \\ &+ \sqrt{(1-C_{ws}^2) \langle s'^2 \rangle \frac{2dt}{T_0}} \eta_i(t) + C_{ws} \sqrt{\langle s'^2 \rangle \frac{2dt}{T_0}} \xi_i(t) \end{aligned}$$

$$R_i^{2'}(t+dt) = R_i^{2'}(t) + 2A_3s_i'dt$$

$$C_{ws} = \langle w's' \rangle / (v_{rms} \sqrt{\langle s'^2 \rangle})$$

 T_0

velocity/supersaturation auto-correlation

Large eddy turn over time

$$\begin{aligned} 1-D \ stochastic \ model/2\\ \frac{d\langle s'R^{2'}\rangle}{dt} &= A_1\langle w'R^{2'}\rangle + 2A_3\langle s'^2\rangle - \frac{\langle s'R^{2'}\rangle}{\langle \tau_s\rangle}\\ \frac{d\langle w'R^{2'}\rangle}{dt} &= 2A_3\langle w's'\rangle - \frac{\langle w'R^{2'}\rangle}{T_0}\\ \frac{d\langle s'^2\rangle}{dt} &= 2A_1\langle w's'\rangle - 2\frac{\langle s'^2\rangle}{\langle \tau_s\rangle}\\ \frac{d\langle w's'\rangle}{dt} &= A_1v_{rms}^2 - \frac{\langle w's'\rangle}{\langle \tau_s\rangle} \quad \text{Steady state} \Rightarrow\\ \langle s'^2\rangle_{qs} &= A_1^2v_{rms}^2\langle \tau_s\rangle^2\\ \langle s'R^{2'}\rangle_{qs} &= 2A_3A_1^2v_{rms}^2\langle \tau_s\rangle^2T_0 = 2A_3\langle s'^2\rangle_{qs}T_0\\ \text{and consequently} \end{aligned}$$

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$$\sigma_{R^2} = \sqrt{8}A_3A_1v_{rms}\langle\tau_s\rangle(T_0t)^{1/2} = \sqrt{8\langle s'^2\rangle_{qs}}A_3(T_0t)^{1/2}$$

since $v_{rms} \simeq Re_{\lambda}^{1/2}v_{\eta}$ and $T_0 \simeq 0.06Re_{\lambda}^{1/2}\tau_{\eta}$

$$\sigma_{R^2} \simeq 0.7 A_3 A_1 \nu^{1/2} \langle \tau_s \rangle Re_\lambda t^{1/2}$$

Standard deviation does not depend on dissipation and so on small scales but is proportional to scale separation

Comparisons with DNS results





- The model approximates the largest simulation
- Smallest simulations are influenced by viscous effects of the smallest scales
- In general the stochastic model tends to overestimate

Error estimation



- DNS/Stochastic model comparison
- Estimate of the supersaturation fluctuations



- Stochastic models overestimates but....
- The error tends to diminishing by increasing the large turbulent scales
- 20% is already a good approximation for evaluating the order of magnitude
- We found an upper limit at no cost

Why these differences?





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Comparison with Large Eddy Simulation



- We want to see the effects of the large scale on droplet condensation
- Maximum cloud size in homogeneous conditions order 100 meters
- Classic Smagorinsky model for the fluid velocity and supersaturation field
- Droplet number: order 10¹⁵ → unfeaseable→ use of renormalization as described in Lanotte et al., 2009
- Parameters $\varepsilon = 10^{-3}$ $m^2 s^{-3}$ $v_{rms} = 0.7$ m/s $Re_{\lambda} = 5000$

Large Eddy Simulation microphysics parametrization



- We assume no sgs-model
- \rightarrow Supersaturation value is not evaluated at the droplet scale
- LES does not evolve the correct number of droplet \rightarrow rescaling
- Equation for droplet radius

$$\frac{dR_i^2}{dt} = 2A_3(s_{res} + s_{sgs}) - - > \frac{d\langle (R_i^{2'})^2 \rangle}{dt} = \frac{d\sigma_{R^2}^2}{dt} = 4A_3(\langle s'R^{2'} \rangle_{res} + \langle s'R^{2'} \rangle_{sgs})$$

- From DNS results $\langle s'R^{2'}
 angle_{res}>>\langle s'R^{2'}
 angle_{sgs}$
- Small scale dynamics is lost \rightarrow underestimation
- Now we have a lower limit at moderate cost

Comparison with large DNS 2048



- DNS 10 days of computations in 4096 cores on the 32nd TOP500 list supercomputer
- LES almost 1 hour in 1 core on my laptop

Model vs LES





- Very good agreement for both standard deviation and correlation
- Model extension for higher Reynolds number
- Significant values of standard variation found after several minutes
- Importance to have longer simulations
- Impact of condensation has been underestimated in the last years

Droplet distribution



RMS is sufficient to characterize pdf

Conclusions and perspectives/1

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- New state of the art simulations able to reproduce condensation growth with time comparable with rain formation
- Standard deviation of square radius fluctuations increases continuously in time according to a power law

Importance of large/small scale separation
Validation of a simple stochastic model that is able to capture all the essential dynamics
Droplet distribution seems to follow a Gaussian curve

Conclusions and perspectives/2

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- Limit of our study: smallest droplets
- Will the rms continue to increase or the final droplet distribution will reach a steady state?
- Combining condensation+collisions
- Difficult but maybe more interesting since
- Condensation—>large scale
- Collision
 Small scales and difficult to include in LES/stochastic models